

Analysis: Based on the data from Figure 2, although the uncertainties are large, a general negative gradient in FeO content can be seen from the north-west to the south-east, where the region mapped as Imbrium ejecta is the highest in Fe (most mafic). This indicates a slightly larger influence from Imbrium than is detectable from the mapping in Figure 1. However, due to the formation of subsequent large craters and mare flooding in the area, it is difficult to determine if this high FeO wt % is due solely to Imbrium ejecta or if there is some other cause.

Although there is also a general negative trend in thorium concentrations heading east, a spike can be seen in the Compton ejecta and in the Bel'kovich region. This is caused by the Compton-Bel'kovich thorium anomaly, marked in Figure 1, which has a very high thorium content relative to the rest of the basin region (see Table 1). This feature has been proposed as the site of the eruption of silicic volcanic rocks (i.e. rhyolite), a very rare phenomenon on the Moon [7].

Other than the high concentrations of FeO in the dark mantle and mare, the other units are all of feldspathic highlands composition, with little variation. Compared with the composition of the other basins previously studied [8-11], Humboldtianum ejecta is low in FeO, even lower than Orientale, which is very iron-poor (Fig. 3). This result suggests that the ejecta from Humboldtianum is extremely feldspathic and the basin crustal target likely consisted of very ancient anorthositic rocks, which appear common in the northern central far side highlands [12].

The TiO₂ content of the ejecta is fairly uniform throughout the region with the two notable exceptions of Compton-Bel'kovich and the terra material (Nt) inside Bel'kovich. The low concentration of TiO₂ inside Bel'kovich could be related to the influence of the thorium anomaly approximately 100 km away.

The average elevation south of the basin is much higher than the average elevation north of it such that the peaks of the northern massifs are approximately equal in elevation as the southern highlands. Moreover, many massifs exist in the northern areas but not in the southern. One possible explanation for this relation is that the basin-forming impact hit at an oblique angle pointing in a northerly direction. It should be noted that the highest elevation in the south is approximately the same as the peaks of the massifs in the north.

Compared with other recently mapped basins, Imbrium, Orientale, and Crisium [9-11], Humboldtianum contains lower levels of Ti and, on average, lower levels of Fe. Notably, impact melts were not found during the mapping process. This is either because they do not exist or they do exist but are now covered by younger impact materials.

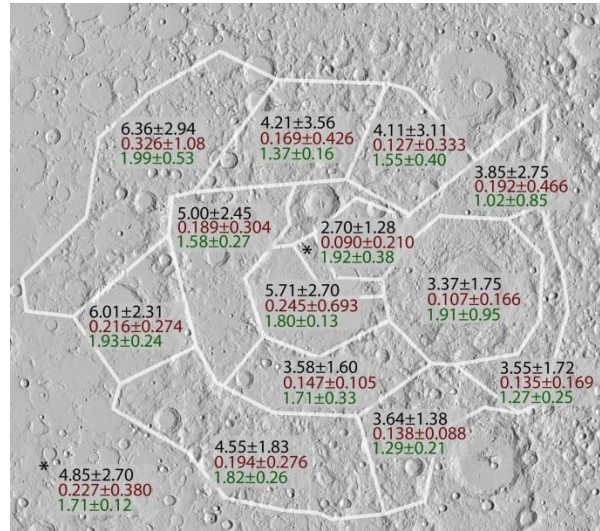


Figure 2 – Regional mineral weight %. Black: FeO, Red: TiO₂, Green: Thorium (ppm). Uncertainties indicate $\pm 1\sigma$

Conclusions: A new geological map of the Humboldtianum basin was created using new data from the LRO, Clementine, and Lunar Prospector missions. The chemical data was overlaid onto the geographic map and related to mapped ejecta units around the basin, and stratigraphic units were recorded and analyzed.

The compositional analysis shows that the different units had similar mineral content. However, there was a noticeable difference in FeO content in the different regions of the basin, a negative content gradient was found to exist from the north-west to the south-east. This indicates that there is a large Imbrium influence on that side of the basin but much less, if any, influence on the side of Humboldtianum basin away from Imbrium.

It was also found that large massifs are found north of the basin while the region to the south is more uniform in elevation, with fewer massifs present in the area.

References: [1] Wilhelms, D. E. (1987) *The Geologic History of the Moon*. USGS Prof. Paper 1348, 302 pp. [2] Lucchitta, B. K. USGS map I-1062. [3] Schultz, P. H., Crawford, D. A. (2016) doi:10.1038/nature18278 [4] Scholten, F. et al. (2012) doi:10.1029/2011je003926 [5] Lucey P.G. et al. (2000) JGR 105, 20297. [6] Lawrence D. et al. (2007) GRL 34, L03201, doi:10.1029/2006GL028530. [7] Jolliff, B. L. et al. (2011) LPSC. [8] Smith, M. C., Spudis, P. D. (2013) LPSC. [9] Martin, D. J. P., Spudis, P. D. (2014) LPSC [10] Murl, J. N., Spudis, P. D. (2015) LPSC [11] Sliz, M. U., Spudis, P. D. (2016) LPSC. [12] Jolliff, B. L. et al. (2000) JGR 105, 4197.